Determination of Total Dissolved Solids in Water Analysis

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THE determination of total solids is one of the oldest determinations in water analysis and has apparently always been taken to represent the amount of material in the water. The difference between the weight of the residue on evaporation of unfiltered and filtered samples of the water has sometimes been taken as a measure of the suspended matter. The loss on ignition has frequently been taken as a measure of the organic matter in the water. Comparison of reported determinations of total solids and loss on ignition with the results of analysis of the actual mineral content indicates, however, that in some analyses the figures for total solids and for loss on ignition have no real significance.

The determination of total dissolved solids is based on the weight of the residue on evaporation after drying at a temperature of 105° to 110° C. (1, 5) or at 180° C. (3). The loss on ignition is determined on the residue used for determining total dissolved solids by careful ignition over a small flame or in a covered metal dish (4). In calculating the quantity of dissolved material, the figure for bicarbonate is divided by 2.03 to account for conversion to carbonate on evaporation, and the quantities of the other constituents as found by analysis are added to give the quantity of dissolved mineral matter. In most water analyses the determined figure for total dissolved solids is greater than the sum of the determined constituents, because of the retention of water of crystallization or water of occlusion and because of small quantities of undetermined constituents. In analyses of certain types of waters, however, the quantity reported for total dissolved solids is less than the sum of the determined constituents because of the volatilization of certain constituents during the evaporation and subsequent drying.

over 5000 made in the Water Resources Laboratory of the U. S. Geological Survey to represent the different types of

CARBONATE WATERS. Natural waters of the most common type are those in which calcium and bicarbonate are the principal radicals. This type is represented by the first two analyses in Table I. The residues from certain of these waters in which magnesium is small, when dried at 180° C., represent fairly accurately the dissolved material, as shown by analysis 1. Many waters of this type, however, contain considerable magnesium, and in some of these waters the determined figure for total solids is less than the sum of the determined constituents, as shown in analysis 2. In over 1000 analyses of waters of this type the determined figure for total dissolved solids was equal to the sum of the determined constituents in 15 per cent, greater than the sum by about 4.1 parts per million in 54 per cent, and less than the sum by about 3.0 parts per million in 31 per cent. The total solids for the whole group of waters averaged 81 parts per million.

The dissolved material in another large group of waters consists chiefly of sodium and bicarbonate (2). This group is represented by analysis 3, which indicates that in a water of this type the determination of total solids, based on the residue dried at 180° C., agrees well with the sum of the determined constituents.

Waters in which sulfate is present in SULFATE WATERS. considerable quantities yield residues on evaporation that contain sulfates of calcium, magnesium, and sodium, with water of crystallization. Most of this water is driven off from the sulfates of sodium and magnesium when the residue is heated at 180° C., but this temperature is not sufficient to dehydrate calcium sulfate completely. Waters of this

TABLE I. MINERAL CONSTITUENTS OF NATURAL WATERS

	SiO ₂	Fe	Са	Mg	Na	K —Parts	HCO ₃	SO:	Cl	NO ₃	TOTAL DISSOLVED SOLIDS AT 180° C.	Sum	FIXED RESIDUE	Color
1	7.9	0.07	52	5.0	2.4	1.7	178	3.9	1.8	0.20	165	163	164	
2	9.0	0.02	32	15	2.6	1.3	171	2.7	2.6	0.53	137	150	133	
3	16	0.07	5.3	3.9	141	3.2	388	5.0	12	1.1	382	378	379	
4	13	0.08	113	29		∽ 3.9	186	407	23	7.3	811	785	774	
5	30	0.12	184	29	12	17	214	48	454	1.0	1076	979	1049	
6	17	0.03	68	19	19	4.7	239	19	29	50	331	343	297	
7	6.3	0.06	17	2.6	1.3	0.8	58	8.2	0.8	0.10	67	66	63	5
8	6.0	0.01	5.2	1.7	5.7	1.4	13	2.3	10	Trace	66	39	51	200

- a Figure obtained by subtracting loss on ignition from figure for total solids.

 - Allan Springs at Price, Ala., 1928. Spring at Village Springs, Ala., 1928. Well, 28 feet deep, at Hendersonville, Pa., 1926. San Juan River near Bluff, Utah, average, 1931.

Natural waters contain varying quantities of silica and the carbonates, sulfates, chlorides, and nitrates of calcium, magnesium, sodium, and potassium. For most waters it is assumed that the silica is present in the colloidal condition and after evaporation is present as SiO₂, although in certain waters, especially some from volcanic regions, the silica in the residue is in combination with the basic elements. Natural waters can be grouped for study according to the predominating acid radicals and in this paper are discussed under four groups-carbonate, sulfate, chloride, and nitrate waters. The analyses in Table I were selected from Well, 182 feet deep, at St. Petersburg, Fla., 1923. Well, 24 feet deep, at Bartonsville, Fla., 1931. Cowpasture River at Clifton Forge, Va., 1930. North Fork of Black Creek, Middleburg, Fla., 1924.

type are represented by analysis 4 in Table I. A great many waters high in sulfate carry enough chloride to affect the relation between the weight of the residues and the amount of dissolved material.

Chloride Waters. Natural waters containing large quantities of chloride also generally contain considerable calcium and magnesium, and the residues from these waters are weighed with difficulty because of the presence of calcium and magnesium chlorides. The determined quantity of dissolved solids in practically all waters of this type is considerably greater than the sum of the determined constitu-

Table II. Loss of Chloride during Evaporation and Heating of Residues from Natural and Artificial Waters

	TOTAL DISSOLVED SOLIDS AT 180° C.		Fe	Ca	Mg	Na+ K	HCO ₈	SO ₄	Cl	NO ₃	Ca	$_{ m Mg}$	Na + K	нсо		Cl	NO3	Loss of Chloride during Evapora- tion and Heat- ing 1 hr. at 180° C.
	Parts per million												-Milligre	am equ	ivaients			Parts per million
1 2 3 4 5 6 7 8	1076 1262 2630 5377 1480	30 24 16 8.4 26	0.12 0.27 0.12 0.36 0.09	184 115 177 254 44 43 37 28	29 61 101 181 40 244 113 192	127 206 643 1418 409 398 340 271	214 154 144 85 259 253 217 174	48 129 253 580 425 1233 353 282	454 504 1342 2662 375 366 546 720	1.0 3.2 1.7 Trace Trace Trace Trace Trace	9.2 5.7 8.8 12.7 2.2 2.1 1.8 1.4	2.4 5.0 8.3 14.9 3.3 20.0 9.3 15.8	5.5 9.0 28.0 61.6 17.8 17.3 14.8 11.8	3.5 2.5 2.4 1.4 4.2 4.1 3.6 2.9	1.0 2.7 5.3 12.1 8.8 25.7 7.4 5.9	12.8 14.2 37.8 75.1 10.6 10.3 14.3 20.3	0.0 0.1 0.0 0.0 0.0 0.0 0.0	50 86 102 86 0 0
	 Well, 162 feet deep, at St. Petersburg, Fla., 1923. Well, 325 feet deep, at Eau Gallie, Fla., 1923. Well, 265 feet deep, at Cocoa, Fla., 1923. Salt Springs near Norwalk, Fla., 1924. 										5. 6. 7, 8.	Natural	l water t l waters	o whic	h solid	magnesi		te was added. esium chloride was

Well, 265 feet deep, at Cocoa, Fla., 1923. Salt Springs near Norwalk, Fla., 1924.

ents. Analysis 5 in Table I represents a water of moderate mineral content of this type.

Natural waters of the chloride type often show loss of chloride on evaporation and during the subsequent heating. The amounts of chloride in eight typical waters and the losses found after evaporation and heating are shown in Table II. In each case the loss of chloride occurred with residues from waters in which the chloride was more than equivalent to the sodium plus potassium. This was also brought out in a series of experiments with natural waters to which sodium chloride, magnesium chloride, and magnesium sulfate were added in varying amounts. The results of such an experiment with natural waters are shown by analyses 6 to 8 in Table II. It was also found that the addition of sodium sulfate or sodium carbonate to a solution of magnesium chloride in such quantities that the sodium becomes more than equivalent to the chloride eliminates the loss of chloride from such a solution.

The loss of chloride is usually not apparent from the weight of the residue on evaporation, because the amount of water held by the residue as water of crystallization is often greater than the loss of chloride.

NITRATE WATERS. Most natural waters have only small quantities of nitrate. In a few, however, the nitrate is present in considerable quantities usually with considerable quantities of other constituents that affect the weight of the residue (analysis 6, Table I). Waters that contain much nitrate may lose as much as 5 or 10 parts per million of nitrate when the residue is left on the steam bath as long as 4 hours. The loss of nitrate is increased during heating at 180° C. for one hour. The losses for seven natural waters are shown in Table III.

TABLE III. LOSS OF NITRATE DURING HEATING OF RESIDUE FROM NATURAL WATERS

(4	Analyses m	ade i	n the '	Wate	r Res	ources	Labor	atory	, U. S	. Geol	ogical S	Survey)
	•										Loss	
											TRATE :	
	bo											ATING
	TOTAL									4	Hours	at
	SOLIDS AT					Na+					on steam	180°
	180° C.	SiO ₂	Fe	Ca	Mg	ĸ	HCO ₂	SO_4	Cl	NO_3	bath	Č.
						per n	illion -					
1	331	17	0.03	68	19	24	239	19	29	50		19
2	406a	11	0.10	88	35	2.24		16	12	42	26	35
3	330a		0.21	72	17	$15^{\dot{a}}$	204	67	14	31	9	18
4	372^{a}			102	18	4.6a		15	19	34	13	30
5	265^{a}		0.05	42	26	12ª	199	16	28	23	7	.7
234567	303ª 209ª	• •	$0.05 \\ 0.04$	70 50	19 15	$\frac{11^a}{3.5^a}$	$\frac{255}{206}$	27 5	$\frac{20}{5.0}$	$\frac{15}{16}$	9 11	13 12
•			0.04	50	10	3,0"	200	J	5.0	10	11	12
	a Calcula				_	_						
						esburg,						
						lboro,			1021			
						ley Spr nite Po						
	5. Well,	100	feet de	en.	at Wi	nitaker	. Va	1931				
	6. Well.	118	feet de	ep,	at Ga	inesbo	ro, Va.	, 193	31.			
	7. Well,	53 fe	et dec	p, a	Gor	esville,	Va., 1	931.				

Colored Waters. Natural waters are often colored by organic matter. Waters of this type are represented by analyses 7 and 8, Table I. The more highly colored water

(analysis 8) showed a much greater loss on ignition than the slightly colored water, and the agreement between the total solids and the sum was better for the water having less color.

DRYING AND IGNITION OF TOTAL SOLIDS. Many natural waters of moderate mineral content give residues on evaporation which when weighed after drying on a steam bath, with no other heating, show a fair agreement between determined total solids and the sum of the determined constituents. For 288 waters whose residues weighed between 100 and 500 mg, the difference between the weight when dry on the steam bath and the weight after heating for an hour at 180° C. was less than 10 mg. on 166 waters, between 10 and 20 mg. on 56 waters, and more than 20 mg. on 66 waters.

More concentrated waters give residues on evaporation which may have to be dried at 180° C. for several hours to obtain a constant weight. For certain types of waters whose residues are hygroscopic and difficult to weigh, it is preferable to evaporate a suitable volume in a weighing bottle that can be stoppered during the weighing. Waters of this type can be evaporated in an open dish rather rapidly and with no appreciable loss through spattering by applying heat from a hot noncorrodible surface above the dish. The residue obtained in this manner gives a better weight for the determination of total solids than that obtained when the water is evaporated in a dish on the steam bath and the residue dried at 180° C. to constant weight, although the loss in chloride is greater when the residue is heated from above. For highly concentrated waters that contain considerable calcium, magnesium, and sodium chlorides, such as brines from oil fields, Zinzalian and Withrow (9) recommend drying the residues at 750° C., making a correction for the amount of calcium and magnesium chlorides in the residue on evaporation. For waters of high concentration Reistle and Lane (6) determined the specific gravity of a sample and used this determination as the basis of the choice of the size of sample to be analyzed. For waters whose specific gravity was greater than 1.025 the total solids were calculated from the determined specific gravity. These authors give a table for the computation of total solids from the specific gravity in a range from 1.020 (total solids, 27,500 parts per million) to 1.140 (total solids, 210,000 parts per million) and state that the total solids computed in this way should check within 2.5 per cent the determined figure for total solids, based on the residue dried at 105° C.

RELATION BETWEEN CONDUCTANCE AND TOTAL SOLIDS. The total salinity of irrigation waters has been estimated from the specific electric conductance of the waters by Scofield and Wilcox (8), who used the method recently described by Scofield (7) for estimating the amount of soluble salts in soils or soil solutions. These authors state:

Where desirable, the total salt content may be estimated as parts per million from the conductance determination by the use of suitable conversion factors. For approximate estimates with irrigation waters having a conductance of 100 or more,

Table IV. Analyses and Specific Conductance of Different Types of Waters (Analyses made in the Water Resources Laboratory, U. S. Geological Survey)

(1224) bed made in the fractional test business, c. o. deciosion but voj)																	
	1											TOTAL ISSOLVED OLIDS AT		Specific Conductance (Reciprocal Ohms × 105			
		SiO ₂	\mathbf{Fe}	Са	Mg	Na		HCO3	SO ₄	Cl	NO3 E	180° C.	Sum	AT 25° C.)	Sum ÷ Conductance		
,						P	arts pe	er milli	on —								
1 2 3 4 5 6 7 8 9 10 11		12 16 27 16 15 8.2 19 6.2 16 8.4	0.05 0.02 0.06 0.10 0.02 0.04 0.04 0.02 0.20 0.08	6.2 8.6 9.4 26 37 108 67 144 70 209 100	2.3 4.0 5.4 14 16 19 9.0 60 44	4.8 4.4 6.5 8.5 17 28 62 16 132 16 217	0.7 1.0 5.6 1.0 4.4 4.8 1.4 2.6 3.2 4.0	28 23 46 94 82 306 160 409 44 380 222	4.3 15 5.3 21 101 25 143 105 446 445 372	4.8 6.0 3.0 7.9 55 76 16 4.5 4.0 231	0.15 1.6 0.20 4.0 1.4 47 2.0 3.8 0.10 0.30	49 60 70 143 241 440 468 539 720 1018 1139	48 62 72 147 236 449 461 526 692 941 1100	6.6 12 11 23 46 69 79 86 105 120 185	7.3 5.2 6.5 6.4 5.5 5.8 6.6 7.8		
12		16	0.10	120	60	129	3.7	195	599	22	15	1129	1061	146	7.3		
Nor Sacr Wel Can	Sacramento River at Sacramento, Calif., April 20, 1932. Northwest Branch at Takoma Park, Md., April 18, 1932. Sacramento River at Sacramento, Calif., April 6, 1932. Well, 704 feet deep, at Middleburg, Va., October 30, 1931. Cance Brook at Irvington, N. J., February 26, 1932. Well, 300 feet deep, at Berryville, Va., October 29, 1931										Dolores River near Cisco, Utah, April 1, 1932. Well, 95 feet deep, at Stephens City, Va., October 28, 1931. Wells at Marion, Ohio, part of public supply, treated water, April 6, 1932. Wells at Marion, Ohio, part of public supply, rew water, April 6, 1932. Colorado River near Grand Junction, Colo., March 11 to 20, 1932. Gunnison River near Grand Junction, Colo., March 11 to 20, 1932.						

- 1. 2. 3. 4.
- Sacramento River at Sacramento, Calif., April 20, 1932. Northwest Branch at Takoma Park, Md., April 18, 1932. Sacramento River at Sacramento, Calif., April 6, 1932. Well, 704 feet deep, at Middleburg, Va., October 30, 1931. Canoe Brook at Irvington, N. J., February 26, 1932. Well, 300 feet deep, at Berryville, Va., October 29, 1931

Residues from waters high in chloride, in which the chloride is more than equivalent to the alkalies, may lose from 50 to 100 parts per million of chloride on heating. Residues from water high in nitrate may lose as much as 30 parts of nitrate on heating.

Waters of high concentration give more consistent results for the total dissolved solids when a small volume is evaporated in a weighing bottle. Time may be saved by evaporation in a dish by application of heat from above.

The specific conductivity of a water may be used to calculate approximately the quantity of dissolved material if the nature of the dissolved material is known.

the specific conductance. This was pointed out by Scofield (7) and is shown by typical results given in Table IV. The factor that must be used to obtain the quantity of dissolved solids from the conductivity of these 12 waters ranges from 5.1 to 7.8 and is not consistent for any range of solids or any type of water. Analyses 10 and 9, which have quite different conversion factors for estimating the solids from the

in which sulfates predominate, the conductance multiplied by

7 will give a figure fairly close to the total salt content in parts per million. If the conductance is much below 100 or if chlorides

predominate in the salt complex, then the conductance multiplied by 6 will be nearer the true value for total salt content.

Natural waters of different types give inconsistent results

when an attempt is made to determine the factor that should

be used to estimate the amount of dissolved material from

conductivity, represent the raw and treated water of a public supply. Analyses 11 and 12, also with quite different factors, represent waters of similar quantities of dissolved solids but different in nature. It is evident that the conductivity is not an accurate measure of the dissolved solids, though it may serve to give approximate results of some practical value.

SUMMARY

The figure for total dissolved solids, based on the weight of the residue on evaporation after heating for 1 hour at 180° C., is reasonably close to the sum of the determined constituents for most natural waters.

Waters of the carbonate type that are high in magnesium may give residues that weigh less than the sum. Natural waters of the sulfate type usually give residues that are too high on account of incomplete drying.

LITERATURE CITED

- (1) Assoc. Official Agr. Chem., Official and Tentative Methods, pp. 35-6 (1916).
- (2) Collins, W. D., and Howard, C. S., Ind. Eng. Chem., 19, 623 (1927).
- (3) Dole, R. B., U. S. Geol. Survey, Water Supply Paper 236, 13-15, 30-1 (1909).
- (4) Drown, T. M., Tech. Quart., 2, 132-5 (1889).
 (5) Mason, W. P., "Examination of Water," 3rd ed., pp. 22-4, Wiley, 1908.
- (6) Reistle, C. E., Jr., and Lane, E. C., U. S. Bur. Mines, Tech. Paper 432, 3-4 (1928).
- Scoffeld, C. S., U. S. Dept. Agr., Circ. 232 (1932).
- (8) Scofield, C. S., and Wilcox, L. V., U. S. Dept. Agr., Tech. Bull. **264**, 21 (1931).
- (9) Zinzalian, George, and Withrow, J. R., Ind. Eng. Chem., Anal. Ed., 4, 210-14 (1932).

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An Inexpensive Flame Shield

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NUMBER of excellent flame shields are available from A the various scientific supply houses at fairly low prices, but even this low cost assumes importance when a relatively large number of students must be supplied. Then, too, in research laboratories, the investigator often needs a flame shield immediately and is inconvenienced if it must be ordered. For these reasons the device described in this article is worth consideration.

The construction of the device is extremely simple. A tin can of the size used for prepared soups is secured, the top cut off smoothly, and a hole punched in the bottom. This hole should be of a size such that it fits tightly the upright tube of the burner, on which the shield is to be used. Adjustment is secured by raising or lowering the can, friction maintaining the position if the hole is of the proper size.

Larger cans may be used if a larger shield be desired. In some cases, gallon cans with both the top and bottom removed and holes punched in the sides for air inlets are advantageous. The relatively short life of the device is compensated by its inexpensiveness.

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